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Description and Genesis of the Western Cold Turkey Creek Field Anomaly, Williston Basin, Bowman County, North Dakota

by

Karyn A. Alme Bachelor of Science, University of North Dakota, 1994

A Thesis

Submitted to the Graduate Faculty

of the

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in partial fulfillment of the requirements

for the degree of

Master of Science

Grand Forks, North Dakota December 2001

This thesis, submitted by Karyn A. Alme in partial fulfillment of the requirements for the Degree of Masters of Science from the University of North Dakota, has been read by the faculty Advisory Committee under whom the work has been done and is hereby approved.

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Ahmad Glancin

This thesis meets the standards for appearance, conforms to the style and format requirements of the Graduate School of the University of North Dakota, and is hereby approved.

Dean of the Graduate School

 $04/01$

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ABSTRACT

The western Cold Turkey Creek structural anomaly is a subsurface structure in Bowman County, North Dakota, near the southern margin of the Williston Basin. The structure was identified by oil explorationists as a possible impact crater, but its origin remained uncertain. The purpose of this study is to attempt to ascertain if sufficient evidence exists to determine that the anomaly is of impact origin.

The study area was confined to a nine township area surrounding the anomalous structure in Bowman County, North Dakota. One hundred six oil exploration wells were identified in the area, and well logs were examined and interpreted for each. The resulting data were used to generate an isopach map for each of twenty-nine lithologic units identified, and an additional five structure contour maps were generated on the tops of conformable units. Units were correlated and two cross-sections through the structure were generated. One core from a well which penetrated the structure was available for study; it was examined and described. Previously published seismic data were also evaluated and interpreted.

Examination of the isopach maps reveals that significant variations in the thickness of multiple lithologic units occur both within the structural anomaly and in the surroun ling study area. Results of this study indicate that all of the lithologic units present within the study area are continuous across the structural anomaly, except one which is discontinuous throughout the study area and doesn't occur in the structure.

Examination of the core revealed nothing suggestive of shock due to explosive impact cratering; only normal sediments were seen. Seismic data and structure contour maps indicate that the shape of the structure is neither circular nor bowl-shaped. Cross sections clearly demonstrate that the anomalous structure developed intermittently from at least the Ordovician until possibly as late as Early Jurassic.

The western Cold Turkey Creek structural anomaly is interpreted to be a complex structure with a long history of development, exhibiting rapid thickness variations over small areas that are suggestive of multiple episodes of up and down motion within the structure. The shape of the structure is inconsistent with that of impact craters, and there is a lack of secondary evidence, such as shatter cones or planar deformation features, to suggest impact as a causal factor. Stratigraphic units appear continuous across the structure and do not show the effects of deformation by hypervelocity impact. The subtle variations in shape, size, and location of the structure suggest that it has developed over a long period of time and therefore, is not the result of an instantaneous event.

INTRODUCTION

General

The western Cold Turkey Creek field anomaly is a subsurface structure located in Bowman County, North Dakota, near the southern margin of the Williston Basin (Fig. 1). The origin of this structure is uncertain however, it exhibits several characteristics similar to other structures in the Williston Basin that have been identified as impact structures.

Purpose

The purpose of this study is to attempt to ascertain whether sufficient evidence exists to determine that the western Cold Turkey Creek field anomaly is of impact origin. Impact events are instantaneous, explosive events. Craters resulting from such events tend to be circular or oval and bowl shaped, with a raised rim, and, depending on size, a central peak. Impact events in the buried rock record arc characterized by abrupt lithology changes, including but not limited to missing units, overturned units, and discontinuous, disrupted or brecciated units. Analysis of the data collected in this study will help to determine whether or not the structure is due to an impact event.

Study Area

The area of study for the purpose of this investigation was confined to a nine township portion of Bowman County, North Dakota: Townships 129, 130, 131 North, Ranges 101, 102, and 103 West. There are several oil fields within these townships,

Figure 1. The position and extent of the Williston Basin, relative to the location of t_{MC} study area. (Modified slightly from Laird, 1956). The star represents the study area location.

either in whole, or in part. Among them is the Cold Turkey Creek field. The western Cold Turkey Creek field anomaly is located in the southwestern comer of Township 130 North, Range 102 West, Bowman County, North Dakota, adjacent to the Cold Turkey Creek oil field (Fig. 2).

Previous Work

From 1969 to 1983, oil explorationists drilled in the anomalous area of the Cold Turkey Creek field due to its seismic patterns bearing a resemblance to those of other impact structures in the Williston Basin (Gerhard et al., 1995). Although there is a lack of published data, explorationists had regarded the Cold Turkey Creek anomaly as a possible impact cratter.

However, Gerhard et al. (1995) demonstrated that small-scale tectonic features could account for the structural anomaly, rather than meteoroid impact or salt dissolution as causal factors. They concluded that the Cold Turkey Creek structure is the result of the interaction between small well-defined basement core blocks moving independently of each other. Gerhard et al. (1995) used orthophotoquad map interpretations of surficial drainage patterns and lineaments, reflection seismic data, and artificial dipmeter reconstruction.

Figure 2. Location of the Study Area, Western Cold Turkey Creek Structural Anomaly, Cold Turkey Creek Field, Bowman County, North Dakota. The star represents the approximate location of the structural anomaly.

REGIONAL SETTING

General

The Williston Basin is a prominent geologic feature in the northern Great Plains. It is a slightly irregular, relatively shallow, intracratonic basin centered in western North Dakota (Carlson and Anderson, 1965). Although the basin includes some 51,000 square miles of North Dakota, it also extends into a significant part of southern Saskatchewan, the southwestern comer of Manitoba, the northwestern portion of South Dakota, and the eastern and north eastern parts of Montana. The outline of the Williston Basin (Fig. 1) is commonly defined by the zero elevation line of the Dakota Sandstone (Laird, 1956) due to its readily identifiable character in wire-line well logs. The Williston Basin is comprised of sedimentary rocks which range in age from late Cambrian to Recent and which reach a maximum thickness of approximately 16,000 feet in western North Dakota (Gerhard et al., 1982). Rocks representing all periods of Phanerozoic time are present within the basin, and occur over a wide area (Gerhard et al., 1982). The relatively quiet tectonic history of the basin and large areal extent of strata facilitate detailed correlation of rock units over large areas (Sawatzky, 1975).

The Williston Basin contains structures of various sizes (Fig. 3). These structures include the large-scale Nesson and Cedar Creek anticlines, other smaller-scale anticlines, faults, and lineaments (Gerhard et al., 1982) as well as impact structures, such as the Red Wing Creek structure (Koeberl and Reimold, 1995) and the Newporte structure (Forsman et al., 1996). The western Cold Turkey Creek Field structural anomaly is also located within the Williston Basin, and is a comparatively small scale structure.

Regional Stratigraphy

The Bighorn Group, of Middle Ordovician to Early Silurian age, contains three formations: the Red River, the Stony Mountain, and the Stonewall (Bluemle et al., 1986). The Red River Formation (Ordovician) is largely composed of limestone and dolomitic limestone. The carbonates of the Red River Formation lie conformably above the terrigenous rocks of the Middle Ordovician Winnipeg Formation (Carroll, 1979). The strata of the Red River Formation are laterally continuous across most of the Williston Basin, and are a source of commercial oil production within the study area (Hvinden, 2001). The Stony Mountain Formation conformably overlies the Red River Formation, and is divided into two members. The lower, or Stoughton member, is mostly interbedded limestone and shale. The upper, or Gunton Member, is composed of dolomite and limestone (Bluemle et al., 1986). The Stonewall Formation lies stratigraphically above the Stony Mountain Formation and is composed of limestone and dolomite with some thin beds of anhydrite (Bluemle et al., 1986), (Fig. 4).

The Interlake Formation (Silurian) is composed of limestone and dolomite, divided into three units: lower, middle, and upper. The lower part is finely crystalline dolomite with thin anhydrite interbedding; the middle portion is fine- to mediumcrystalline dolomite; the upper part is dolomite and limestone, silty and sandy near the base (Bluemle et al., 1986). Although sedimentation was continuous from the beginning of Red River deposition through the Silurian, the top of the Interlake Formation is marked by an unconformity, the result of regressing seas (Gerhard et al., 1982).

Figure 3. Prominent structural features within the Williston Basin of eastern Montana and western North Dakota (Modified slightly from Gerhard et al., 1982).

Figure 4. Generalized stratigraphic column of the study area.

The Ashem Formation (Lower Devonian), is a dark-gray to brick-red dolomite (Bluemle et al., 1936). The study area is located near the southwestern margin of the Ashem Formation's areal extent in the Wiliiston Basin, and as a result, is intermittent to discontinuous within the study area (Carlson and Anderson, 1965).

The Souris River Formation is Middle Devonian. It consists of limestone, dolomite, and thin interbeds of shale, silt and evaporites, representing a cyclic pattern of deposition (Bluemle et al., 1986).

The Late Devonian Jefferson Group is divided into the Duperow and Birdbear Formations. The Duperow Formation lies stratigraphically above the Souris River Formation, and also consists of carbonates, primarily limestone and dolomite. Shale is also present, but in smaller quantities than in the Souris River Formation. The Birdbear Formation is primarily limestone with some dolomite and small amounts of anhydrite (Bluemle et al., 1986).

The Three Forks Formation (Upper Devonian), consists of red and green siltstones and shales, interbedded with dolomite and anhydrite (Carlson, 1967). A thin layer of sandstone occurs at the top of the Three Forks Formation (Bluemle et al., 1986).

The Madison Group ranges in age from Early Mississippian to Late Mississippian, and is divided into three formations: the Lodgepole, the Mission Canyon, and the Charles. The Lodgepole Formation forms the lower portion of the Madison Group. It is a light- to dark-gray, generally dense limestone. The Mission Canyon Formation consists of yellowish-brown to brownish-gray limestone, often fragmental, oolitic, and cherty. It contains intertonguing lenses of anhydrite and shaly dolomitic limestone (Bluemle et al., 1986). The Charles Formation is primarily composed of

evaporites, with interbedded halite anhydrite, mudstone, dolomite, and shale. The Lodgepoie, Mission Canyon, and Charles Formations are facies that have a complex intertonguing relationship with one another, and cut across log markers (Carlson and Anderson, 1967; Bluemle et al., 1986).

The Big Snowy Group is Late Mississippian and is divided into two formations: the Kibbey and the Otter. The Kibbey Formation conformably overlies the Charles Formation and is divided into three lithologic units. The lowest of these units is the Kibbey shale, which is reddish in color, silty and interbedded with gypsum. The middle unit is the Kibbey limestone. It varies in color from white to brown and is often dense and dolomitic. The "Kibbey lime" is an excellent log marker. The upper unit is the Kibbey sandstone which is reddish to light-gray in color and fine- to medium-grained (Bluemle et al., 1986). The Otter Formation lies strati graphically above the Kibbey sandstone and occurs predominantly as variegated greenish- and reddish-gray carbonaceous shale. Some thinly bedded fossiliferous limestone occurs within the Otter as well.

The Minnelusa Group lies unconfonnably on top of the Big Snowy Group. The Pennsylvanian Tyler Formation is the lowermost lithologic unit of the Group. It consists of shale and limestone. The remainder of the Minnelusa Group is early Permian to Late Pennsylvanian and is largely composed of pinkish-gray sandstone and dolomite (Bluemle et al., 1986).

The rocks of the Opeche Formation are Permian, and consist of orange-red, slightly dolomitic shale and siltstone. Streaks of anhydrite and gypsum are also present (Carlson, 1967).

The Minnekahta Formation lies stratigraphically above the Opeche Formation and is Permian. It is composed of a creamy pink and purple mottled limestone (Bluemle et al., 1986).

The Spearfish Formation lies conformably atop the Minnekahta Formation. It has three members; the lowermost stratigraphically is the Belfield Member. The Belfield consists of grey shale and red siltstone. The second member of the Spearfish Formation, which lies above the Belfield Member, is the Pine Salt. The third member is the Saude, which is composed of siitstone, sandstone, and interbeds of shale (Carlson and Anderson, 1965). The Belfield and the Pine Salt are Permian, while the Permian-Triassic boundary occurs somewhere within the Saude (Bluemle et al., 1986).

Jurassic strata lie unconformably above the Spearfish Formation and represent a primarily terrigenous sequence of deposition in the Williston Basin. Jurassic strata in the Williston Basin reach a maximum thickness of 1,200 feet (367 meters) in extreme northwestern North Dakota, and extend over all but the easternmost part of the state. Jurassic rocks generally consist of limestone, shale, gypsum and anhydrites.

The Dakota Group (Lower Cretaceous) is an interbedded sandstone and shale interval. The Inyan Kara Formation is the lowest unit of the group, consisting of lightgray sandstone and shale. The Mowry is the uppermost formation of the Dakota Group. It is a shale unit, medium- to dark-gray in color with tra \sim s of light-blue-gray bentonitic clay. The top of the Mowry Formation is marked by a radioactive zone, and thus is a distinctive log marker (Bluemle et al., 1986).

The Colorado Group (Middle Cretaceous) includes the Greenhorn and Niobrara Formations. The Greenhorn Formation is a dark-gray calcareous shale with thin

interbeds of shaly limestone. It is also a good log marker, having distinctive electric and radioactive characteristics. The Niobrara Formation is the uppermost unit of the Colorado Group. It is characterized as a medium-light-gray to medium-gray calcareous shale (Biuemle et al., 1986).

The Pierre Formation (Upper Cretaceous) is a thick shale unit. It varies in color from light- to medium- or dark gray, and is generally noncalcareous. The Pierre shale is commonly fissile and may be flaky to blocky. Deposition of the Pierre shale occurred in a marine offshore setting (Biuemle et al., 1986).

METHODS

Introduction

The research for this study involved several steps. A literature search was conducted to obtain reference material about the Cold Turkey Creek field, the Williston Basin, its stratigraphy, structures and tectonic history, as well as impact crater morphology, structure, and occurrence. Depth and thickness data were recorded from well logs from wells in and around the structural anomaly. This data was used to generate isopach and structure contour maps. Core samples from a well which penetrates the structure were observed, and seismic maps of the structural anomaly were also evaluated.

Well Logs

Since the structure in question is a subsurface structure, data were obtained about it remotely. One hundred six wells (Fig. 5) were identified in nine townships around and containing the Cold Turkey Creek oil field, where the structure is located (Appendix A). The North Dakota Geological Survey and the North Dakota Industrial Commission: Oil and Gas Division maintain records of oil exploration wells drilled within the state of North Dakota. These logs arc available in paper forms, microfilm, and microfiche. All three formats were utilized to obtain data for this study. At least two types of well logs are normally recorded for each well: dual laterologs and compensated neutron logs. Dual

R. 103 W. R. 107 W. R. 101 W.

Figure 5. Location of wells from which well logs were examined during this study, and the location of the well from which core was examined. Township boundaries are shown for scale; townships are six miles on each side. Filled circles represent producing wells; open circles represent dry wells; the square represents the well with core samples

laterologs, which typically record gamma ray, spontaneous potential and resistivity data about the strata, were used for the purpose of this investigation. Compensated neutron logs, which record gamma ray and porosity index data, were not used, as porosity values were unnecessary to the context of this investigation. Wire-line well logs from all one hundred six wells in the area were examined, and interpreted; log signatures were lentified and their depths recorded for each of twenty-nine different lithologic units. ranging in age from Ordovician to Cretaceous. These units, from oldest to youngest, are: Red River Formation, Stony Mountain Formation, Stonewall Formation, Interlake Formation, Ashem Formation, Souris River Formation, Duperow Formation, Birdbear Formation, Three Forks Formation, Lodgepole Formation, Mission Canyon Formation, Charles Formation, Kibbey shale, Kibbey limestone, Kibbey sandstone, Otter Formation, Tyler Formation, Minnelusa Group, Opeche Formation, Minnekahta Formation, Spearfish Formation: Belfieid Member, Pine Salt Member, Saude Member, Jurassic strata, Inyan Kara Formation, Mowry Formation, Greenhorn Formation, Niobrara Formation, and the Pierre Fonnation.

The depth and thickness data obtained from the well logs were compiled into a database (Appendix B) using a spreadsheet for ease of analysis. Depths to lithologic unit tops were calculated by subtracting the formation top picks from the kelly bushing to compute the new depth relative to sea level. These data were contoured by SURFER 7 ® (Golden Software, 1999) to generate a structural contour map of each lithologic unit surface. Maps are a projection of the state plane onto a flat surface. The thickness of each lithologic unit was calculated by subtracting the depth of the top of the underlying

unit from the depth to the top of the unit being measured. These thickness values were then contoured by SURFER 7 \degree (Golden Software, 1999) to generate isopach maps of each lithologic unit, except the Red River, for which no underlying unit was identified and thus no thickness values were calculated. Contour maps were generated by SURFER 7 ® (Golden Software, 1999). The method used for gridding the data were point kriging which used a linear variogram. with slane of one, and anisotropy of zero. A grid of 200 by 200 lines was used, and maximum and minimum values for X and Y were contained within the nine township study area. Contours were smoothed by SURFER 7[°]. The search radius used all data, and the number of points varied for each map generated, ranging from 90 to 98 data points, with the exception of the isopach map of the Pierre Formation, which used 37 data points. Data for lithologic unit tops were also used to construct stratigraphic columns which were then correlated between wells in order to generate two cross-sectional diagrams through the structure.

Seismic Structural Maps

Recent seismic data for the area in question were still proprietary at the time of this writing and were unavailable for study. Therefore, previously published seismic data (Gerhard et al., 1995) were used for interpretation.

Core

One core from the structure was available for study. The core from NDGS well number 4654, International Nuclear Corp. #1 - 61 John M. Susa et al, was examined to determine general lithologic characteristics, and identify any possible definitive evidence of impact. The core is maintained by the North Dakota Geological Survey at the Wilson

M. Laird Core and Sample Library on the University of North Dakota Campus, Grand Forks, North Dakota. The total length of the observed core was 103 feet (31 meters), and was neither complete nor continuous. The core contained sections from the Mission Canyon Formation and the Red River Formation (Appendix C).

DISCUSSION

Impact cratering is a fundamental geologic process, the evidence for which can be seen on the surface of most rocky bodies in the solar system. More than one hundred fifty impact craters have been identified on Earth (Koeberl and Sharpton, 2001), and several have been identified within the Williston Basin (Sawatzky, 1977; Gerlach et al., 1995; Koeberl et al, 1996). An impact crater is the result of the transfer of the kinetic energy of a projectile to the surface of a planet. Crater features are affected by the velocity and mass of the projectile, the nature and composition of the target materials, and the manner in which the target materials respond to the impact-induced shock wave and subsequent excavation of materials (Hamblin and Christiansen, 1990, p. 423). Morphologic characteristics of impact craters are a function of the energy expended during the explosive impact event. A simple crater has a bowl-like depression, with smooth floors and walls and ejected debris surrounding the raised crater rim. Large, complex impact craters, due to higher energy events, are characterized by central peaks or rings, crater rim terraces, and possibly outer rings. An impact is an instantaneous event, and is reflected as such in the rock record.

Most impact craters on Earth have been modified by erosion, subsequent sedimentation, and the dynamic processes of the lithosphere (Hamblin and Christiansen, 1990, p. 199). Nonetheless, certain diagnostic characteristics help to define and identify impact craters in the rock record. Impact craters are typically round or oval. The target

material at the site is shocked and excavated by the explosion, causing ejection,

liisuc transport, and re-deposition of target material. Therefore, the stratigraphy of the target rocks is highly disturbed; it is marked by missing beds (material excavated during impact), disrupted and overturned beds, and brecciated material. Target rocks are subjected to pressures of many hundreds of GPa and temperatures exceeding several thousand degrees Celsius. As a result, minerals undergo structural and phase changes uniquely characteristic of the extreme high pressures and strain rates of hypervelocity meteorite impact (Koeberl et al., 1996). Several microscopic shock metamorphic effects have been recognized as indicators of impact. Planar deformation features (PDFs) are microscopic features within mineral grains. They consist of narrow planes of glassy material arranged in parallel sets having distinctive orientations with respect to the grain's structure (Koeberl and Sharpton, 2001). High pressure polymorphs of quartz; coesite and stishovite, are also produced by the high pressures associated with impact and have been recognized as indicators of an impact event. Shatter cones are another indicator of impact. They are megascopic cone-shaped structures seen in rocks near impact craters. They have an appearance of regular thin grooves which radiate from an apex. They range in size from less than one centimeter, to more than a meter (Koeberl et al, 1996). Shatter cones form under shock pressures in the range of two to thirty GPa.

Impact structures have been identified within the Williston Basin (Fig. 3). The Newporte structure (Ordovician) is a petroliferous, subsurface impact crater located just south of the North Dakota border with Saskatchewan. Approximately 2 miles (3.2 km) in diameter, it is a simple crater, having a circular depression surrounded by a raised rim.

PDFs were identified by Gerlach et al, (1995) in thin sections of rocks from wells in the structure's rim. Formations are discontinuous across the structure. The Viewfield crater in Saskatchewan is also a subsurface, bowl-shaped crater with a raised rim. It is considered to be of Jurassic - Triassic age and is identified by seismic data and overturned strata at its rim (Sawatzky, 1975). The Red Wing Creek structure is a larger and more complex impact crater. It too is a petroliferous, subsurface structure, and is located in the center of the Williston Basin, in western North Dakota. Koeberl and Reirnold (1996) confirmed the impact origin for the Red Wing Creek structure by identifying PDFs which indicate a shock pressure of at least twelve GPa in samples of well cuttings from the center of the structure. The Red Wing Creek structure has a central uplift, approximately 9.6 miles (15.4 km) in diameter, surrounded by an approximately 3.3 mile (5.3 km) wide ring depression, which is in turn surrounded by a raised rim of deformed rocks. Drilling in the central uplift encountered breccia and shatter cone fragments (Brenan et al., 1975).

RESULTS

Red River Formation

As the base of the Red River Formation was not identified on well logs, no isopach map of the Red River Formation was constructed. However, the structural contour map of the top of the Red River indicates a gradual downward slope toward the northeast portion of the study area. A very slight depression is apparent at the western Cold Turkey Creek structure $\{Fig. 6\}$.

Stony Mountain Formation

The Stony Mountain Formation, which conformably overlies the Red River Formation, reaches a maximum thickness of 192 feet (59 meters) near the southwestern comer of the study area, where two structures exhibit thickening trends within the State Line oil field (Fig. 7). The Stony Mountain Formation gradually thins toward the northeast comer of the study area to a minimum thickness of 81 feet (25 meters). The unit thickens slightly at the anomalous structure west of the Cold Turkey Creek field.

Stonewall Formation

The Stonewall Formation maintains an almost constant thickness of 110 feet (34) meters) throughout the study area. However, an anomalous area occurs in the southwest comer of the State Line Field, where the Stonewall Formation thickens to 140 feet (43 meters). Another anomalous area coincides with the minimum thickness (65 feet, 20

Figure 6. Structure contour map on the Top of the Red River Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey line represents generalized oil field boundaries; refer to Figure 2 for field names. Datum is mean sea level. Contour interval $= 25$ feet.

Figure 7. Isopach map of the Stony Mountain Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey line represents generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 10 feet.

meters) of the formation in the northern part of the Amor field in the western part of the study area (Fig. 8).

Interlake Formation

The Interlake Formation (Silurian), which conformably overlies the Stonewall Formation, has a maximum thickness ot 410 feet (125 meters) in the eastern part of the study area. The Interlake formation gradually thins westward, reaching a minimum thickness of 217 feet (66 meters) at the extreme western part of the study area, in the southwes^{*} portion of the Amor field (Fig. 9). A shallow, irregularly shaped depression occurs in the northeastern portion of the Cold Turkey Creek field. The thickness of the Interlake Formation exhibits very little change at the structural anomaly west of the Cold Turkey Creek field, and the unit is continuous across the structure.

Ashem Formation

Stratigraphically above the Interlake Formation is the Ashem Formation, which forms the base of the Manitoba Group. The Ashem Formation is absent over much of the study area. A few thin lenses occur at the extreme west and southwest edges of the study area, and also in the area north of the Amor field and the northwestern most portion of the Cold Turkey Creek field. The maximum thickness (19 feet, 6 meters) occurs in the southeast quarter of T. 130 N., R. 102 W (Fig. 10). The Ashem Formation is not present within the Cold Turkey Creek structural anomaly.

Souris River Formation

The Soui is River Formation is generally greater than 40 feet (12 meters) thick over the study area, with a slight thickening trend on the western-most edge of the study area, near the Amor field. The maximum thickness of 144 feet (44 meters), occurs

Figure 8. Isopach map of the Stonewall Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 10 feet.

Figure 9. Isopach map of the Interlake Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

Figure 10. **Isopach** map of the Ashern Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval $= 5$ feet.

within the Cold Turkey Creek oil field in the central portion of T. 130 N., R. 102 W. The minimum thickness of 8 feet (2.4 m) occurs within the southeastern portion of the Amor field (Fig. 11). The thickness of the Souris River Formation is slightly greater on the northern side of the structural anomaly than the southern side, and the unit is continuous across the structure.

Duperow Formation

The Duperow Formation exhibits a slight thickening trend from the southwest to the northeast across the study area, reaching a maximum thickness of 220 feet (67 meters) in the northeast comer of the study area. The minimum thickness of 134 feet (41 meters) occurs as an anomalous area in the center of the study area (Fig. 12), within the Cold Turkey Creek field. From the structure contour map, it is apparent that the Duperow Formation dips downward into the Williston Basin, in a northeasterly direction. A shallow circular depression is present at the structural anomaly west of the Cold Turkey Creek field (Fig. 13).

Birdbear Formation

The isopach map of the Birdbear Formation indicates that the thickness is generally less than 100 feet (30 meters). There is a general thickening trend toward the eastern portion of the study area, with an anomalous area in the Amor field, where rapid thickness changes are evident (Fig. 14). The thickness of the Birdbear Formation is almost constant across the structural anomaly.

Three Forks Formation

The Three Forks Formation generally thickens to 61 feet (19 meters) in the

Figure 11. Isopach map of the Souris River Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

Figure 12. Isopach map of the Duperow Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 10 feet.

Figure 13. Structure contour map on the Top of the Duperow Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figue 2 for field names. Datum is mean sea level. Contoui interval = 25 feet.

Figure 14 Isopach map of the Birdbear Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field locations; refer to Figure 2 for field names. Contour interval $= 10$ feet.

eastern portion of the study area and thins to zero feet at the extreme western portion of the study area. The thickness of the Three Forks Formation at the western Cold Turkey Creek structure does not appear anomalous (Fig. 15), and the unit is continuous across the structure.

Lodgepole Formation

The Lodgepole Formation is generally greater than 540 feet (164 meters) thick in the western portion of the study area and generally less than 580 feet (177 meters) in the northern part (Fig. 16). Fairly rapid thickness changes occur where the Lodgepole Formation reaches a maximum thickness of 646 feet (197 meters) in the center of the study area, just to the south of the western Cold Turkey Creek structural anomaly, and the Formation reaches a minimum thickness of 447 feet (136 meters) within the structural anomaly.

Mission Canyon Formation

Rapid thickness changes are again apparent in the Mission Canyon Formation, where maximum thickness (524 feet, 160 meters) and minimum thickness (180 feet, 55 meters) for the study area occur within the western Cold Turkey Creek structural anomaly. The thickness of the area surrounding the structure is fairly constant, with a slight thickening trend toward the east (Fig. 17).

Charles Formation

The Charles Formation maintains a fairly constant thickness across the study area of between 450 and 490 feet (137 and 149 meters, respectively). However, rapid

Figure 15. Isopach map of the Three Forks Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval $= 10$ feet.

Figure 16. Isopach map of the Lodgepole Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

Figure 17. lsopach map of the Mission Canyon Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

thickness changes occur at the western Cold Turkey Creek structural anomaly, where a maximum thickness of 943 feet (287 meters) occurs. East of the structural anomaly, the minimum thickness of 296 feet (90 meters) is also accompanied by rapid thickness changes over ϵ small area (Fig. 18). The area of maximum thickness is evident on the structural contour map as well, where it stands out as a steep slope from a local high to an adjacent local low. The area of minimum thickness is less distinctive on the structural contour map, although it is still apparent as an area of decreased slope. The surrounding study area exhibits a general slope downward to the north-northeast, in the direction of th< center of the Williston Basin (Fig. 19).

Kibbey shale

The Kibbey shale, the oldest unit of the Kibbey Formation, generally ranges in thickness from 65 feet (20 meters) to 85 feet (26 meters) across the study area (Fig. 20). Rapid thickness changes are notable in the structural anomaly west of the Cold Turkey Creek field, where the Kibbey shale reaches a maximum thickness of 132 feet (40 meters). The minimum thickness of 37 feet (11 meters) occurs in the center of the southern-most part of the study area.

Kibbey limestone

The Kibbey limestone is the middle unit of the Kibbey Formation, and is the thinnest unit apart from the Ashem (which is absent over much of the study area). The Kibbey limestone generally ranges in thickness across the study area from 11 feet (3.5) meters) to 36 feet (11 meters). The maximum thickness of 83 feet (11 meters) appears as an anomalously thick area in the extreme western portion of the study area, in the

Figure 18. Isopach map of the Charles Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval $= 25$ feet.

Figure 19. Structure contour map on the top of the Madison Group: the Charles Fonnation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Datum is mean sea level. Contour interval = 25 feet.

Figure 20. Isopach map of the Kibbey Formation shale. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 10 feet.

northern part of the Amor field (Fig. 21). The unit is continuous across the western Cold Turkey Creek structural anomaly.

Kibbey sandstone

The Kibbey sandstone is the uppermost unit of the Kibbey Formation. It has a variable thickness across the study area, ranging from a minimum of 19 feet (5.8 meters) to a maximum of 178 feet (54 meters). Both the maximum and minimum thicknesses occur in the eastern portion of the Cold Turkey Creek field, in rather close proximity to one another. Several wells in the Amor field at the extreme western side of the study area indicate slightly greater than average thickness changes among them (Fig. 22). Another area of greater than average thickness occurs in T. 129 N., R. 102 W., just north of the Ash field

Otter Formation

The Otter Formation ranges in thickness from a minimum of 16 feet (5 meters) to a maximum of 227 feet (69 meters) and is variable throughout the study area. Although the thickness of the Otter Formation changes slightly across the western Cold Turkey Creek structural anomaly, the unit is continuous (Fig. 23).

Tyler Formation

The isopach map of the Ty'er Formation (Fig. 24), indicates that the thickness of the unit across the study area is slightly variable, with a minimum thickness of 62 feet (19 meters) at the extreme western edge of the study area. The maximum thickness of 400 feet (122 meters) occurs within the western Cold Turkey Creek structural anomaly, and is identified by the tight concentric contours on the isopach map. Although there is a

Figure 21. Isopach map of the Kibbey Formation limestone. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 10 feet.

Figure 23. Isopach map of the Otter Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 20 feet.

R. 103 W. R. 102 W. R. 101 W.

Figure 24. Isopach map of the Tyler Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

substantial thickness change at the western Cold Turkey Creek structural anomaly, the Tyler Formation is continuous across the structural anomaly.

Minnelusa Group: Broom Creek - Amsden Interval

The Minnelusa Group: Broom Creek - Amsden interval reaches a maximum thickness of 322 feet (98 meters) in the extreme eastern part of the study area,whiie the minimum thickness of 118 feet (36 meters) occurs near the center of the study area, within the Cold Turkey Creek field. The structural anomaly is not marked by rapid thickness changes, and the Minnelusa Group is continuous across the structure (Fig. 25).

Opeche Formation

The Opeche Formation reaches a minimum thickness of 44 feet, (13 meters) in the northwest part of the study area, near the northern portion of the Amor field. The maximum thickness (166 feet, 50 meters) is located in the center of the study area, within the Cold Turkey Creek field. Although the thickness of the Opeche Formation changes slightly across the anomalous structure west of the Cold Turkey Creek field, the unit is continuous across the structure (Fig. 26).

Minnekahta Formation

The thickness of the Minnekahta Formation is relatively constant across the study area, ranging from 39 feet (12 meters) to 59 feet (18 meters). The western Cold Turkey Creek structural anomaly is rather inconspicuous, as the unit's thickness changes very little across the structure (Fig. 27). However, the structure contour map of the Minnekahta Formation depicts a slightly elevated area centered on the western Cold

Figure 25. Isopach map of the Minnelusa Group: Broom Creek - Amsden interval. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

Figure 26. Isopach map of the Opeche Foimation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

Figure 27. Isopach map of the Minnekahta Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 5 feet.

Turkey Creek structural anomaly. Generally, the Minnekahta Formation trends northward, dipping downslope toward the center of the Williston Basin (Fig. 28).

Belfield Member

The oldest member of the Spearfish Formation is the Belfield Member. The isopach map of the Belfield Member illustrates that there is little variation in the thickness of the Ξ elfield Member across the study area, which ranges from 113 feet (34) meters) to 202 feet $(6!$ meters). There is a slight thinning trend toward the west and northwest and toward the southeast while a slight thickening occurs along a southwest to northeast diagonal through the center cf the study area. While there is a slight thickness change across the western Cold Turkey Creek structural anomaly, it does not appear to be significant, and the unit is continuous across the structure (Fig. 29).

Pine Salt Member

The Pine Salt Member is the middle member of the Spearfish Formation, and thickens from east to west-northwest across the study area. The Pine Salt is absent in the southeast corner of the study area, and gradually thickens westward, where a maximum thickness of 278 feet (85 meters) occurs in the northern Amor field (Fig. 30). A slight thickening trend occurs at the structural anomaly. The Pine Salt is continuous across the structural anomaly west of the Cold Turkey Creek field.

Saude Member

The Saude Member is the uppermost and youngest member of the Spearfish Formation, and has a variable thickness across the study area. The Saude Member reaches a minimum thickness of 53 feet (16 meters) in the Ash field at the extreme

Figure 28. Structure contour map on the Top of the Minnekahta Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Datum is mean sea level. Contour interval = 25 feet.

Figure 29. Isopach map of the Belfield Member of the Spearfish Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 10 feet.

Figure 30. Isopach map of the Pine Salt Member of the Spear fish Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

southern part of the study area, and reaches a maximum thickness of 314 feet (96 meters) in the eastern part of the study area. The thickness of the Saude changes slightly across the western Cold Turkey Creek structural anomaly, nonetheless the unit is continuous (Fig. 31).

Swift - Piper Interval

The Jurassic Swift - Piper Interval ranges in thickness from 343 feet (104 meters) in the eastern part of the study area, to 647 feet (197 meters) across the study area in the extreme western part of the Amor field. The Jurassic strata are continuous and the thickness is fairly constant across the western Cold Turkey Creek structural anomaly (Fig. 32).

Inyan Kara Formation

The Inyan Kara Formation reaches its maximum thickness of 662 feet (202 meters) near the center of the study area, in the eastern portion of the Cold Turkey Creek field. The minimum thickness of 454 feet (138 meters) occurs in the northeastern part of the study area. Thickness changes across the western Cold Turkey Creek structural anomaly are minimal and the structure is not distinguishable on the isopach map of the Inyan Kara Formation (Fig. 33).

Mowry - Skull Creek interval

The Mowry - Skull Creek interval (Cretaceous) has a minimum thickness of 365 feet (111 meters) in the southeast comer of the study area, thickening gradually toward the center of the Cold Turkey Creek field, then thinning slightly northwest of the Cold Turkey Creek field, before thickening again to reach a maximum thickness of 418 feet

R. 103 W. R. 102 W. R. 101 W.

Figure 31. Isopach map of the Saude Member of the Spearfish Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

R. 103 W. R. 102 W. R. 101 W.

Figure 32. Isopach map of the Jurassic Swift - Piper Interval. Township boundaries are shown for scale; townships are six miles on each side. Points represent weil locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

Figure 33. Isopach map of the Inyan Kara Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

(127 meters) in the northwestern most comer of the study area. The Mowry - Skull Creek interval is continuous across the western Cold Turkey Creek structural anomaly, where the thickness is distinguished by a very slight thinning of the surrounding strata (Fig. 34). The structure contour map of the Mowry Formation illustrates a general trend dipping to the east and northeast. Structure contours highlight a slightly elevated area at the western Cold Turkey Creek structural anomaly (Fig. 35).

Greenhorn - Belle Fourche Interval

The Greenhorn - Belle Fourche interval ranges in thickness from a minimum of 428 feet (130 meters) in the southeast central part of the study area to a maximum thickness of 506 feet (154 meters) in the extreme western part of the study area. There is a slight thinning trend eastward across the western Cold Turkey Creek structural anomaly, though the unit is continuous (Fig. 36).

Niobrara - Carlile Interval

The Niobrara - Carlile interval has a fairly constant thickness throughout the study area, with a minimum thickness of 597 feet (182 meters) in the extreme northwestern comer and a maximum thickness of 729 feet (222 meters) in the southwest corner of the study area. The western Cold Turkey Creek structural anomaly is not distinguished by thickness changes: the Niobrara - Carlile interval maintains a fairly constant thickness across the structure (Fig. 37).

Pierre Formation

The Pierre Formation is the most recent unit examined in the study area; it is also the thickest. The Pierre reaches a maximum thickness of 2319 feet (707 meters) in the

Figure 34. Isopach map of the Mowry - Skull Creek Interval. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 10 feet.

Figure 35. Structure contour map on the Top of the Mowry Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Datum is mean sea level. Contour interval $= 25$ feet.

Figure 36. Isopach map of the Greenhorn - Belle Fourche Interval. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 10 feet.

Figure 37. Isopach map of the Niobrara - Carlile Interval. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval = 25 feet.

northwestern most comer of the study area and a minimum thickness of 1481 feet (451 meters) in the southern portion of the study area. As only 37 of 107 well logs yielded data about the Pierre, the validity of this map is somewhat questionable. Although the western Cold Turkey Creek structural anomaly does not appear distinguishable from its surroundings on this map (Fig. 38), it should be noted that insufficient data has been collected to make a description of Pierre thickness or structure in that location.

Cross Sections

Six wells in and around the structure were used to generate each cross-sectional diagram. These diagrams depict the thicknesses of each unit and the depth to the top of each unit. North Dakota Geological Survey wells numbered 4654, 9317, and 5459 are drilled into the structural anomaly west of the Cold Turkey Creek field. Wells numbered 5823 and 6398 are located north of the structure, and wells 5865 and 8021 are located south of the structure. Wells numbered 11255 and 9963 are located west of the structure, and well number 8982 is located to the east (Fig. 39).

At the base of each well, is the top of the Red River Formation (Figs. 40 and 41). There is some variation in the depth to the top of the Red River Formation. The thickness of the Stony Mountain Formation is fairly consistent, although the slope is slightly variable. The same observations can be made for the Stonewall, and Interlake Fonnations. The thickness of the Souris River Formation is slightly more variable than that of the underlying units, and it is notable that the Ashem Formation is present only in NDGS well number 11255, the western-most well of the west - east cross-section (Fig. 41). The Ashern Formation is a discontinuous unit across the study area, and is not

Figure 38. Isopach map of the Pierre Formation. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed grey lines represent generalized oil field boundaries; refer to Figure 2 for field names. Contour interval $= 100$ feet.

R. 103 W. R. 102 W.

Figure 39. Location of cross-sections. Township boundaries are shown for scale; townships are six miles on each side. Points represent well locations. Dashed line represents north - south cross-section. Solid line represents west east cross-section.

Figure 40. A north to south cross-section through the western Cold Turkey Creek field anomaly, illustrating thickness variations of lithologic units. Vertical exaggeration = $10x$.

Figure 41. A west to east cross-section through the western Cold Turkey Creek field anomaly, illustrating thickness variations of lithologic units. Vertical exaggeration = 10x.

present within the western Cold Turkey Creek structural anomaly or in the wells immediately nearby. The thicknesses of the Duperow, Birdbear, and Three Forks Formations are also fairly consistent across the structure in both the north - south and west – east cross-sectional views. These formations are continuous across the structure and are nearly horizontal, lacking in any dramatic thickness changes. A very slight depression in these formations is evident in well 5459.

The Lodgepole Formation is continuous across the structure, despite thickness changes. The Lodgepole Formation thins across the structure, notably in wells 9317 and 5459, which are located near the approximate center of the structure and just east of center, respectively (Fig. 16)

The Mission Canyon, middle formation of the Madison Group, appears thinner in wells 4654 and 9317, within the structural anomaly, than in wells surrounding the structure and is also anomalously thicker than nearby wells in well 5459. The unit can be correlated between wells across the structure, and is considered continuous, despite thickness changes. The relative depths to the surface of the Mission Canyon and subsequent decreased slope from the north and west toward NDGS well 4654, suggest the presence of a depression to the northwest of well 4654 on the northwest side of the structure. This indicates a shift in the location of the depression on the Mission Canyon top, relative to the location of the depressed area on the top of the underlying Lodgepole Formation.

The Charles Formation is the uppermost unit of the Madison Group, and is continuous across the western Cold Turkey Creek structure. The unit thins at NDGS well

number 4654 on the northwest side of the structure, and then thickens at NDGS well number 9317. While the Charles Formation is only slightly thicker at well number 5459 than it is east and west of the area surrounding the structure, it is noticeably thicker than in wells north and south of the structure. The elevation at which the top of the Charles Formation occurs in NDGS well number 4654 is lower than its elevation in surrounding wells, suggesting the presence of a depression on the northwest side of the structure. The top of the Charles Formation occurs at a higher elevation in wells 9317 and 5459, indicating the presence of a local high on the center and south side of the structure.

The Kibbey shale overlies the Charles Formation, and is a substantially thinner unit. The Kibbey shale is continuous across the structure, though it too varies in thickness; the unit thickens at well 9317. The thickness is fairly constant in the other wells. The relative depths to the top of the Kibbey shale indicate that it occurs at a higher level in wells 9317 and 5459 than it does in neighboring wells, and it occurs as a local low in well number 4654.

The Kibbey limestone, middle unit of the Kibbey Formation, is also continuous across the structure. The unit gradually thickens south and eastward to well 5459, before thinning slightly in well number 8982 which is adjacent to and due east of the structure. The relative depths to the top of the Kibbey limestone indicate that in wells 9317 and 5459, the elevation of the Kibbey lime is higher than the surrounding area, and that it occurs at a lower elevation in well 4654 than either the surrounding area or the structure itself.

The Kibbey sandstone is continuous across the structure and gradually thickens north and westward, with a slightly greater thickness occurring in NDGS well number 4654. The top of the Kibbey sandstone appears elevated and horizontal in an east – west direction between wells 9317 and 5459 within the structure. To the north and west, the surface of the Kibbey sandstone slopes downward to a local low elevation at well 4654, on the northwest side of the western Cold Turkey Creek structural anomaly.

The Otter Formation is continuous across the structure, although the thickness is variable, both within the structure and in surrounding areas. The lop of the Otter Formation mimics the surface of the underlying Kibbey Formation, with the surface of the Otter Formation occurring at a higher elevation in well 5459 than in surrounding wells. The surface is also elevated above the surrounding terrain in well 9317, but less so than in well 5459. There is a thickening trend in well 5459 and a thinning trend in well 9317, while the thickness of the Otter Formation is only slightly greater in well 4654 than it is in wells surrounding the structure.

The Tyler Formation unconformably overlies the Otter Formation, and is continuous across the structure, although its thickness is not constant. North Dakota Geological Survey well number 4654 indicates a pronounced increase in thickness, more than double the thickness present in adjacent wells. The relative depths to the top of the Tyler Formation indicate that the elevation of the surface of the Tyler is elevated at well 5459 above the rest of the stiucture. The surface of the Tyler Fomation dips to the northwest and is accompanied by a raised area at well 5459.

The top of the Minnelusa Group exhibits a slight depression to the northwest of the structure, with a slightly raised area occurring at well 5459. The unit generally dips very slightly to the northwest. The thickness of the Minnelusa Group is less variable than underlying units, with a slight thickening trend to the south and east apparent in both cross-sections (Figs. 40 and 41). The unit is continuous across the structure.

The Opeche Formation exhibits a slight thinning trend to the south and the east, though the unit is continuous across the structure. Although the Opeche Fonnation unconformably overlies the Minnelusa Group, the topography of the top of the Opeche Formation is very similar to that of the underlying unit. There is a slight depression to the north and west of the structure, accompanied by a slightly raised area at weii 5459.

The Minnekahta Formation lies stratigraphically above the Opeche Formation, and is continuous across the western Cold Turkey Creek structure. The thickness of the Minnekahta Formation exhibits very little variation in *he cross-sections. The top of the Minnekahta Formation exhibits a nearly horizontal slope, dipping slightly westward in the west – east cross-section (Fig. 41). Again, a slightly raised area occurs in well 5459 and also in well 9317.

Overlying the Minnekahta Formation is the Bel field Member, the lowermost member of the Spearfish Formation. The thickness of the Belfield Member shows little variation in the cross-sections, and is continuous across the structure. The west – east cross-section illustrates that the unit dips slightly to the northwest. However, the surface of the unit appears to be more variable in the north - south cross-sectional view, where

the Belfield Member occurs at a slightly higher elevation in well 5459 than in surrounding areas.

The Pine Salt Member of the Spearfish Formation is continuous across the western Cold Turkey Creek structural ar maly. Although its thickness is slightly variable within the structure, the Pine Salt thins in well 4654, and then thickens in wells to the north and west. Other wells to the south and east exhibit little variation in thickness. The top of the Pine Salt Member, although nearly horizontal, exhibits a very slightly raised area at wells 5459 and 9317.

The uppermost member of the Spearfish Formation is the Saude Member, which is continuous across the structure, even though its thickness is only slightly variable. The top of the Saude Member is nearly horizontal, with only the westernmost well having a higher elevation than the others in the cross-section.

The Jurassic and Cretaceous strata above the Triassic Saude Member are all continuous units across the western Cold Turkey Creek structural anomaly. All of these units have a nearly horizontal orientation, with a slight northeast dip direction. None of these units exhibit significant thickness changes at or near the structure. The thicknesses of the Jurassic and Cretaceous strata show little variation anywhere in either of the crosssections (Figs. 40 and 41).

Seismic Structural Maps

Two seismic structural maps of the western Cold Turkey Creek structural anomaly were available in a previously published work (Gerhard et al, 1995).

The seismic structural contour map of the top of the Ordovician Winnipeg horizon utilizes a contour interval of 3 msec, which is similar to, but not the same as measurements of feet. The contour interval of msec represents the amount of time taken for the signal to travel down to the surface before it is reflected back to the seismograph. Since the time traveled is proportional to the velocity, it yields an approximation for distance to the reflecting unit. The Winnipeg Formation lies stratigraphically below the Red River Formation, and thus is representative of data not available in the well log data base collected for this study. Figure 42 illustrates the western Cold Turkey Creek structural anomaly as a rhombic anti formal structure located in the southwestern comer of Township 130 North, Range 102 West, in sections 29, 30, 31, and 32. North Dakota Geological Survey well numbered 4654 is located on the northwest flank of the structure, while the well numbered 9317 is positioned near the highest area of the structure in section 30, just south of the approximate center. North Dakota Geological Survey well numbered 5459 is due east of well number 9317, and is located in the downward sloping area adjacent to the flat, elevated top, in section 29. The eastern side of the structure is marked by a very steep slope, while the western side slopes much more gradually to the surrounding terrain. In the northeast comer of section 29 is a depression. Elsewhere on the map, the structure contours indicate that the depth increases in the north and northwest portions of the map, indicating a gradual slope dipping to the north and northeast toward the center of the Williston Basin.

The seismic structural contour map on the top of the approximate Mississippian Mission Canyon Formation illustrates a more complicated structure (Fig. 43). The

Figure 42. Seismic structural contour map on the top of the Ordovician Winnipeg horizon (Modified slightly from Gerhard et al., 1995).

Figure 43. Seismic structural map on the surface of the approximate Mississippian Mission Canyon horizon (Modified slightly from Gerhard et al., 1995).

the depth indicated. Generally, greater depths are indicated by structure contours in the north and northeastern portions of the seismic structural contour map of the Mission Canyon, suggesting a general north-northeast dip direction for the Mission Canyon Formation (Fig. 43) surrounding the anomalous structure. The southwest side of the structure has a steep slope, which ends in an irregularly-shaped shallow depression, which trends northwest to southeast, extending approximately from well 4654 to south of well 9317 and eastward to south of 5459. Adjacent to and sharply contrasting the depression is an elevated knob, just northeast of well 9317, in the southeast comer of section 30. Section 29 is dominated by the presence of a second elongated, irregular depression trending northwest to southeast. Between the two depressions is a northwest to southeast trending antiform.

Core

The core from North Dakota Geological Survey well number 4654 located on the northwestern part of the western Cold Turkey Creek structural anomaly in Township 130 North, Range 102 West, Section 30, SWNE, was examined and described (Appendix C). 103 feet of discontinuous core from well 4654 contained sections from the Mission Canyon Formation and the Red River Formation. Core intervals are in feet below the surface.

The core interval from 7568 feet to 7642 feet, although discontinuous, contained 71 feet (21 meters) from the Mississippian Mission Canyon Formation. This predominantly consisted of limestone, massive to finely laminated, with carbonate mud and anhydrite nodules throughout (Appendix C). This description is consistent with

published descriptions of typical Mission Canyon characteristics common throughout the Williston Basin. Nothing atypical was observed: no unusual structures, textures, or unexpected lithologies.

The core interval from 9440 feet to 9567 feet was also discontinuous, and represented only a portion (33 feet, 10 meters) of the Red River Formation from well 4654. The core in this interval was composed of limestone, mottled gray, finely laminated to massive, with occasional fossils and occasional stylolites (Appendix C). This description is consistent with other published descriptions of typical Red River Formation rocks elsewhere in the Williston Basin.

INTERPRETATION

In general, several characteristics are used to identify an impact crater in the rock record, namely: shape, structure, stratigraphic signature, and secondary indicators, such as PDF's, high pressure polymorphs of quartz, and shatter cones. Additionally, an impact crater is the result of an instantaneous event. In order to evaluate the origin of the western Cold Turkey Creek structural anomaly as a possible result of impact, the anomaly's structure and stratigraphy were examined and described to determine if the structure represented an instantaneous geologic event, later simply buried by additional sedimentation. If the structure could be determined to be the result of impact, data collected by this study should be able to determine the geologic time for that impact.

The results of this study indicate that the structure does not represent an instantaneous geologic event. The earliest development 0f the structure is at least as old as the Winnipeg Group (Ordovician), if not older. The seismic data indicates the presence of the structure within the Winnipeg, and 'acking any information about strata older than this, the possibility exists that the structure may have developed earlier. Analysis of the data obtained from this study suggests that the western Cold Turkey Creek structural anomaly may have begun, in part, as a rhombic, antiformal structure on the surface of the Ordovician Winnipeg Group.

On the surface of the Red River Formation, the structure appears as an anomalous elevated area on an otherwise smooth slope into the center of the Wilhston Basin. Well

log data for the Red River, Stony Mountain, and Stonewall Formations above the Winnipeg, indicate that the strata are nearly horizontal in the area of the structure, and that there is little variation in the thickness of these units. Very slight thickness changes are apparent in wells 4654 and 9317 in the Interlake and Souris River Formations, suggesting that there was some localized and subtle relative motion up and down during the deposition of these units. The Duperow, Birdbear, and Three Forks Formations above the Souris River maintain almost constant thicknesses within the three wells which penetrate the structure. The surface and slope of these units reflect the thickness changes in the underlying Interlake and Souris River Formations, and otherwise appear relatively unaffected by the structural anomaly.

The entire Madison Group, however, is marked by changes in the area of the structural anomaly. Rapid thickness changes over short distances again indicate relative up and down motion within each unit and between wells. The surface of the Lodgepoie Formation is irregular, and the thickness of the unit changes sounatically between wells 4654, 9317, and 5459 (Figs 16 and 41). In the Mission Canyon Formation, the surface is also irregular, but it does not mimic that of the underlying Lodgepole Formation. The seismic structural contour map of the approximate Mississippian Mission Canyon Formation indicates how complex the structure has become since deposition of the Ordovician Winnipeg Group. The rhombic antiformal structure has become a northwestto southeast- trending antiform, accompanied by depressions on its northeast and southwest flanks (Fig. 43). The depression to the northeast occurs in the same location as a depression northeast of the rhombic antiformal structure in the Winnipeg (Fig. 42). The structure is not circular. 1 he thickness of the Mission Canyon Formation is highly

variable across the structure, with a dramatic increase in thickness at well 5459 (Figs. 17 and 41). Further, the structure continues to change through the deposition of the Charles Formation, as indicated by additional dramatic thickness changes between wells within the structure.

Superimposed on top of the Madison Group, the Kibbey and Otter Formations reflect the complicated underlying structure. The Kibbey shale exhibits dramatic thickness changes over a small area, especially between wells 5654 and 9317, where the difference is nearly 50 feet (Fig. 20). This is a substantial change for a relatively thin unit, and indicates that the structure continued to develop during deposition of the Kibbey shale. The Kibbey limestone, by contrast, maintains an almost constant thickness between the three wells in the structure (Fig. 21), suggesting a hiatus in structural development. Although the surface of the Kibbey limestone reflects the underlying structure, it remains unaltered by the process that caused the thickness changes in the Kibbey sandstone, stratigraphically above. The Kibbey sandstone like the Kibbey shale, exhibits significant thickness changes between the three wells which penetrate the structure. This indicates renewed development of the structural anomaly.

The Otter Formation also reflects the underlying structure, but is further complicated by a substantial increase in thickness at well 5459. At this location, the unit has a greater thickness than elsewhere in the structure, as well as the surrounding area. This suggests relative up and down motion between wells 9317 and 5459, though more pronounced than such motion in the Kibbey sandstone.

Stratigraphically above the Otter Formation, the top of the Tyler Formation exhibits a decrease in slope across the structure, becoming somewhat more horizontal.

However, evidence for the continued development of the structure is indicated by the substantial gain in thickness at well 4654, and also by the thin nature of the unit at well 5459 (Figs. 24 and 41). Strata above the Tyler Formation exhibit very little expression of the structure below, though the structure is still faintly apparent as a slight rise above the surrounding Tyler terrain.

In both the Minnelusa and Opeche Formations, thickness changes are apparent between wells within the structure, and those outside the structure. Stratigraphically above, the thicknesses of the Minnekahta Formation and Belfield Member of the Spearfish Formation are nearly constant across the structure. However, the overlying Pine Salt has a slightly more variable thickness across the structure. It should be noted that the Pine Salt has a variable thickness across the entire study area, and therefore it is difficult to determine whether thickness variations within the structure are due to structural development.

The Saude Member of the Spearfish Formation exhibits some differences in thickness between wells 4654, 9317, and 5459, although they are not as dramatic as those of underlying units. Jurassic and Cretaceous strata above the Saude maintain almost constant thicknesses across the structure, indicating that further development of the structure had ceased sometime during or after the deposition of the Saude, perhaps during the Jurassic.

An impact crater, if present, should be marked by: highly disturbed beds, missing beds, and overturned beds near the crater rim. Examination of isopach maps and crosssectional diagrams indicate that although thickness changes occur through time at the western Cold Turkey Creek structural anomaly, all units are present and continuous

across the structure. The only exception is the Ashem Formation, which is absent at the structural anomaly and discontinuous across the study area, as the study area is located at the erosional edge of the Ashem's extent in the Williston Basin. There is no evidence to suggest overturned or otherwise disrupted strata.

The core from North Dakota Geological Survey well number 4654 was found to contain predominantly limestones, typical of Mission Canyon and Red River rocks. The core description corresponds to descriptions of type sections of Mission Canyon and Red River rocks elsewhere in the Williston Basin. There was no apparent quartz or feldspar which would suggest the presence of coesite, stishovite or planar deformation features (PDFs). Shatter cones were not observed in either interval, nor were any other extraordinary deformation features observed.

CONCLUSIONS

1) Ail of the lithologic units examined in this study are continuous across the western Cold Turkey Creek structural anomaly, with the exception of the Ashem Formation, which is discontinuous throughout the surrounding area because the study area is located on the erosional margin of the unit. The Ashem does not appear to pinch out at the structural anomaly. Further, the lithologic units do not appear to have been overturned or disturbed by compressive and expansive pressures associated with a hypervelocity impact.

2) The seismic data indicates that the shape of the structure changes through time. The structure appears as a rhomic antiform in the Winnipeg (Ordovician), and as an elongated antiform accompanied by irregularly shaped depressions in the Mississippian Mission Canyon. The structure contour maps, although lacking somewhat in well control, do not suggest the presence of a single circular crater, but rather a structure which changes shape through time. The structure does not exhibit a circular or oval shape, with a smooth floor and sides, or a raised rim.

3.) There are no secondary indicators of impact. The core examined for this study does not suggest the presence of any shock features in the Mission Canyon or Red River intervals. Rather, the rocks from the core section are typical of representative rock sections elsewhere in tue basin.

4.) Examination of the isopach maps of each stratigraphic unit and analysis of the cross-sections indicate that there is not a single episode of excavation followed by infilling. Rather, there are multiple, complex episodes of both gradational and rapid thickening and thinning of units, not in a single location, but in several locations both within the structure, and throughout the study area. The structure is not the result of an instantaneous event, but rather developed through a succession of multiple events over time.

The Western Cold Turkey Creek structural anomaly is a complex structure. A lack of seismic data and minimal well control prohibit the interpretation of a detailed structural reconstruction. However, data obtained and analyzed during the course of this study clearly indicate that the structure has developed over time. The structure is not the result of an instantaneous event, and thus, by definition, is not an impact crater. Further, the shape of the structure is inconsistent with that of impact craters, there is a lack of evidence for secondary indicators of an impact, and the units in the study area appear to be continuous across the structure; there are no missing, discontinuous, or highly disturbed beds that would suggest the presence of an impact crater.

Through the course of this study, several other anoma'ous structures in addition to the western Cold Turkey Creek field anomaly were observed within the study area. Data obtained from this study could be used to generate cross-sections of these other anomalous structures in an attempt to better understand the structural history of both the western Cold Turkey Creek field anomaly and the study area at large. Interpretation of the western Cold Turkey Creek field anomaly and other anomalous structures in the study

area associated with oil fields may help to clarify our understanding of other small-scale features within the Williston Basin.

APPENDICES

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APPENDIX A

Names and Location of Wells. All wells from which well logs were studied are located in Bowman County, North Dakota, and are listed according to their North Dakota Geological Survey well number. The locations of these wells are described as follows: Townships and ranges are north and west respectively. (T = township, $R = range$, $S =$ section, $Q =$ quarter of the quarter section description, and $KB =$ kelly bushing: elevation above sea level). The operator and well name are from records of the North Dakota Geological Survey.

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APPENDIX B

Group, Formation and Member Tops. Wells are listed according to their North Dakota Geological Survey well number. All wells are located in Bowman County, North Dakota. Depths are indicated in feet from Kelly Bushing; Pierre Formation, Niobrara Formation, Greenhorn Formation, Mowry Formation, Inyan Kara Formation, Spearfish Formation / Saude Member, Pine Salt Member, Belfield Member, Minnekahta Formation, Opeche Formation, Minnelusa Group, Tyler Formation, Otter Formation, Kibbey Formation / Kibbey sandstone, Kibbey limestone, Kibbey shale, Charles Formation, Mission Canyon Formation, Lodgepolc Formation, Three Forks Formation, Birdbear Formation, Dupcrow Formation, Souris River Formation, Ashem Formation, Interlake Formation, Stonewall Formation, Stony Mountain Formation, Red River Formation. (no data = missing data; not picked = insufficient information to determine elevation; not here = lithology is absent in well log).

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APPENDIX C

Core Descriptions. Core samples are from North Dakota Geological Survey well number 4654, Operator: International Nuclear, Name: John M. Susa et al #1-61. Depths are from the Kelly Bushing as given on the core boxes. Core descriptions were made using a hand lens and were classified according to Folk (1959, orbonates), Picard (1971, mudrocks), and Maiklem et al. (1969, evaporites). Rock colors were based on the Goddard et al. (1963) Rock Color Chart.

121 NDGS Well No. 4654 International Nuclear, J. M. Susa #1-61 SWNE, Sec. 30, T. 130 N., R. 102 W. Bowman County

Description

Depth *(*feet)

Mission Canyon Formation 7568 -7571 Limestone, light brown to tan, finely laminated with occasional stylolites, effervescent, occasional fossils occur throughout, calcareous. At 7568.5 large stylolitic boundary is present. 7571 -7571.5 Limestone, light brown to tan, finely laminated with occasional stylolites, becoming more stylolitic downward than above, effervescent, occasional fossils occur throughout, calcareous. 7571.5 - 7572 Limestone, medium grey, stylolitic, calcareous. 7572-7573.5 Limestone, light brown to tan, finely bedded, interbedded with fossiliferous limestone. 7573.5-7574 Limestone, grey, massive with occasional laminations. 7574- 7575.5 Limestone, grey, forms matrix with anhydrite nodules. 7575.5 - 7576 Limestone, fragmented, grey matrix with angular tan limestone particles scattered throughout, particle sizes range from 1 mm to 2 cm. 7576 - 7678 Limestone, grey brown, massive to finely laminated, calcareous, occasional stylolites. 7578 - 7579.5 Limestone, light brown and grey interbedded limestone, laminated and stylolitic. 7579.5 -7581.5 Limestone, light brown and grey, finely laminated, grading downward to massive limestone with fine grey laminations. 7581.5 -7583 Limestone, dark grey, alternating layers with medium grey limestone, finely laminated, occasional fossiliferous layers occur, having a thickness not greater than 1 cm. 7583 - 7586 Limestone, grey, very finely laminated, interbedded with carbonate mud, occasional stylolites occur near the bottom.

Anhydrite lenses occur near base.

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